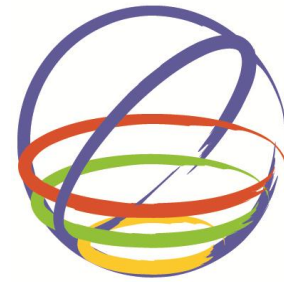


A Body-Fitted Coordinates Based Method for Simulation Earthquake with Inhomogeneous Feature in Space



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SUMMARY:

Spatial variation of earthquake ground motion is an important phenomenon that cannot be ignored in the design and safety of strategic structures. However, almost all the procedures for the evaluation of variation are assumed that the random field is homogeneous in space. It obvious that reality does not fully conform to the assumption (such as a bridge is located in different soil condition). A Body-Fitted Coordinates based method is proposed to estimate and describe of the local spatial variations for the earthquake ground motion; it need not to make the assumption that the random field of earthquake is homogeneous in space. An analysis of spatial variability of seismic motion in smart-1 array monitored in Lotung, Taiwan demonstrates this methodology.

Key words: earthquakes; spatial correlation; dynamic time warping distance; Multidimensional scaling; seismograph array

1. INTRODUCTION

Spatial variation of earthquake ground motion is an important phenomenon that cannot be ignored in the design and safety of strategic structures (Zerva A. et al., 1994). The seismic motion, which could be considered as the result of the complex wave propagation through a heterogeneous soil, is affected by many factors such as propagation path, source mechanism and the amplification effect of the subsurface layers et al. Thus, the record from the seismograph array is not the simple duplication of the traveling wave with the shift of time, but the difference will be amplified when the distance between stations become larger. How to make the description and estimation of correlation structure of spatial ground motion in efficient and robust way is a fundamental prerequisite for problems of simulation of earthquake ground motion and design of structures with spatially extended foundations and lifeline systems.

The development of procedure for evaluate the spatial instationarity has lagged well behind the development of procedure for evaluate the spatial stationary (or isotropy/homogeneous in space). The classical way to describe the spatial variation of earthquake ground motion problem is performed by using the correlation or coherence function (Abrahamson. et al., 1991; Harichandran. 1986), however:

- 1) It is assumed that the random field is homogeneous in space, it is obvious that reality does not fully conform to that assumption (eg, one bridge on rock and the other on alluvium).
- 2) The amplitude variability of seismic ground motion ought to be incorporated into the description of spatial variation of seismic motion, However, the correlation or coherence function only can reflect the phase variation (Zerva A. et al., 1994);

Earthquake as a spatial-temporal field, the focus of the researchers are usually on the temporal non-stationary (Yinfeng Dong. et al., 2010), The lessons learned during the devastating earthquakes of the last decade tell us: just as the character of temporal non-stationary has big influence on seismic response of the structure, the influence of spatial non-stationary on some spatially extended structure

also cannot be ignored. In the design of structures, the reasonable input should take the feature of spatial non-stationary of earthquake into account. It is well known: nearly every significant discovery of feature of earthquake will bring huge breakthrough of seismic design of structure (Zerva A. et al., 2009), the research of spatial non-stationary is of great significance.

In this study, a new method is proposed to describe the spatial variation of earthquake ground motion. This methodology was motivated by problems how to investigate and visualize the variations of earthquake ground motion in space if the field is inhomogeneous.

The fruit of our research can be applied at least two aspects: (1). To help design "optimal" networks of sampling stations for observing these distributions; (2) To investigate the feature of earthquake and as a fundamental prerequisite for the problems of producing properly correlated motions in inhomogeneous space to help the design of the long-span structure.

This article explains our method, with descriptions of the application of the three principal tools: dynamic time warping distance, multi-dimensional scaling and thin-plate spline interpolation and we also present an application of the method to characterize the local spatial correlation of the ground motions which have been recorded obtained from closely spaced arrays (smart-1 arrays event-40).

2. OUTLINE OF THE METHOD

Our approach requires two tools: the Multidimensional Scaling (MDS) and the technique of spatial interpolation. The former generates a low dimensional space (two or three) representation visualizing proximities of the sampling stations; the latter provides smooth mappings of the geographic representation of the sampling stations into their MDS representation.

2.1 Multidimensional Scaling technique

2.1.1 The dynamic time warping distance

For the accelerogram record a_{ij} in n station (sample point), each has N independent realizations. The accelerogram record can be given in a $(n \times N)$ data matrix,

$$A = \begin{bmatrix} a_{11} & a_{12} & \cdots & a_{1N} \\ a_{21} & a_{22} & \cdots & a_{2N} \\ \vdots & \vdots & \ddots & \vdots \\ a_{n1} & a_{n2} & \cdots & a_{nN} \end{bmatrix}$$

The coherence correlation function or coherency coefficient is often used as a similarity measure of earthquake ground motion (Smith, S. et al., 1979), the (Pearson) coherency function is defined as:

$$r_{i,j} = \frac{S_{ij}}{(S_{ii}S_{jj})^{1/2}} \quad (1)$$

and

$$S_{ij} = \frac{1}{N-1} \left[\sum_{k=1}^N a_{ik} a_{jk} - \frac{1}{N} \left(\sum_{k=1}^N a_{ik} \right) \left(\sum_{k=1}^N a_{jk} \right) \right] \quad (2)$$

but some researchers (Todorovska et al., 1997) have found the simple and crude definitions of coherence implies high degree of variability of strong motion even for short separation; It may be from the limitations of coherency function or correlation coefficient itself. The correlation coefficient requires that two series have the same length, the values of two series have point-to-point correspondence, and the weight of each pair of differences is equal. Due to such a correspondence, it may not be suitable to be applied to the similarity measurement of complex series with shift and stretching of amplitude.

The dynamic time warping distance (DTW) proposed by Berndt and Cliffrd (1996) is designated to depict the greatest similarity between series by calculating the minimum distance between them, which is defined as follows:

Let $A(a_1, a_2, \dots, a_n)$ and $B(b_1, b_2, \dots, b_m)$ be two series with the length of n and m , respectively, and an $n \times m$ matrix M can be defined to represent the point-to-point correspondence relationship between A and B , where the element M_{ij} indicates the distance $d(a_i, b_j)$ between a_i and b_j . Then the point-to-point alignment and matching relationship between X and Y can be represented by a time warping path:

$W = \langle w_1, w_2, \dots, w_K \rangle$, $\max(m, n) \leq K < m + n - 1$, where the element $w_k = (i, j)$ indicates the alignment and matching relationship between a_i and b_j . If a path is the lowest cost path between two series, the corresponding dynamic time warping distance is required to meet

$$DTW(A, B) = \min_W \left\{ \sum_{k=1}^K d_k, W = \langle w_1, w_2, \dots, w_K \rangle \right\} \quad (3)$$

$d_k = d(a_i, b_j)$ is the Minkowski distance of the k element of the path W and is denoted as follow:

$$d(A, B) = \left(\sum_{i=1}^n |a_i - b_j|^p \right)^{1/p} \quad (4)$$

When $p = 2$ the distance between two series is called Euclidean Distance.

Then the formal definition of dynamic time warping distance between two series is described as

$$\begin{aligned} DTW(\langle \rangle, \langle \rangle) &= 0 \\ DTW(A, \langle \rangle) &= DTW(\langle \rangle, B) = \infty \\ DTW(A, B) &= d(a_i, b_j) + \min \begin{cases} DTW(A, B[2:-]) \\ DTW(A[2:-], B) \\ DTW(A[2:-], B[2:-]) \end{cases} \end{aligned} \quad (5)$$

where $\langle \rangle$ indicates empty series, $[2:-]$ indicates a sub-array whose elements include the second element to the final element in an one-dimension array, $d(a_i, b_j)$ indicates the distance between points a_i and b_j . The DTW distance of two-time series can be calculated by the dynamic programming method based on accumulated distance matrix, whose algorithm mainly is to construct an accumulated distance matrix:

$$\begin{aligned} r(i, j) &= \min \{ r(i-1, j), r(i, j-1), r(i-1, j-1) \} \\ &+ d(a_i, b_j) \end{aligned} \quad (6)$$

Any element $r(i, j)$ in the accumulated matrix indicates the dynamic time warping distance between series $A_{1:i}$ and $B_{1:j}$. Series with high similar complexity can be effectively identified because the best alignment and matching relationship between two series is defined by the dynamic time distance.

2.1.2 Optimal scaling

Multidimensional Scaling (MDS) is applied as a statistical technique to visualize dissimilarity data in this section.

Let $\Delta = [\delta_{ij}]$ and $D = [d_{ij}]$, are two $N \times N$ matrices indexed by i and j , where the proximity or data

value connecting object i with object j is represented by Δ , d_{ij} refers to the Euclidean space (composition map) between objects i and j . The main objective of MDS is to represent these dissimilarities as distances between points in a low dimensional space (or called composition) such that the distances d_{ij} correspond as closely as possible to the dissimilarities δ_{ij} .

Classical Metric Multidimensional Scaling is a basic form of MDS, Classical MDS is called metric methods because the relationship between δ_{ij} and d_{ij} depend on the numerical or metric properties of the dissimilarities. It only works under the assumption that the geometrical model fits the data perfectly. However, it is often not possible to construct an explicit functional form D such that the mapped dissimilarities D of an empirical data set match sufficiently well metric distances. Therefore, we would like to assign numerical values to the optimal approximations of the transformed proximities to the distances in the geometrical representation. These numerical values are usually called disparities, and they are denoted by \hat{d} and $d_{ij} \approx \hat{d}_{ij} = f(\delta_{ij})$; it should be restricted by the monotonic constraint: $p_{ij} \leq p_{lk} \Rightarrow D(p_{ij}) \leq D(p_{lk})$;

The coordinates in the distance function (in composition map) and the function f which allows transforming the proximities into distances are estimated by minimizing the following badness of fit function (usually called *stressor* S -function in the context of MDS) or called optimal scaling. The stress-function is given by:

$$S = \left(\frac{\sum \sum (\hat{d}_{ij}^2 - d_{ij}^2)^2}{\sum \sum d_{ij}^4} \right)^{1/2} \quad (7)$$

Optimal scaling aims to find a transformation of the data that fits as well as possible the distances in the MDS solution, and find a matrix \hat{D} by some iterative algorithms to make the S minimum. $(x_{1i}, x_{2i}, \dots, x_{pi})$, where p is the dimensional of the final composition map ($p=1, 2, \dots, n$ ect).

The method is summarized as follow:

1. Select the initial matrix $(x_{1i}^0, x_{2i}^0, \dots, x_{pi}^0)$;
2. Calculate the distance matrix $D_k = [d_{ij}^0]$, and elements of Matrix D_k , $d_{ij}^0 = \|x_i, x_j\|$ is the Euclidean space;
3. From $\Delta = [\delta_{ij}]$ and $D_k = [d_{ij}^0]$ use a monotonous regression Pool-Adjacent Violators (PAV) (Ayer et al, 1995) to get $\hat{D}_k = [\hat{d}_{ij}^k]$;
4. For the arbitrary i, j, k, l if $d_{ij}^k < d_{lk}^k \Rightarrow \hat{d}_{ij}^k < \hat{d}_{lk}^k$ then $d_{ij}^k = \hat{d}_{ij}^k$, otherwise $\hat{d}_{ij}^k = \hat{d}_{lk}^k = (d_{ij}^k + d_{lk}^k) / 2$;
5. Calculate the new coordinate of node:
$$x_{qi}^{k+1} = x_{qi}^k + \frac{\theta}{n-1} \sum_{j \neq i} \left(1 - \frac{\hat{d}_{ij}^k}{d_{ij}^k}\right) (x_j^k - x_i^k); \quad (8)$$

n, θ are respectively the number of the node the and iterative step.
6. Use the coordinate of node in step 5 to calculate the Euclidean d_{ij} .
7. Calculate the S according to formula (7).
8. If $S < \varepsilon$, the calculation is end, otherwise return to step 3.

Finally, the result $(x_{1i}, x_{2i}, \dots, x_{pi})_i$ ($i = 1, 2$) is the solution of MDS; If we chose $p=2$, the coordinate can be express as (x_{1i}, x_{2i}) . We refer to the plane of geographical (y_{1i}, y_{2i}) of the sampling stations as G plane, and we refer to the plane of Composition coordinates (x_{1i}, x_{2i}) as C plane.

2.2 The boy-fit grid based depiction of mapping

Classical research of investigating the variation of earthquake in space is on the assume that the ground motion is homogeneous in space (or the function to describe the variation of earthquake is the functions of the separation distance between stations, but independent of absolute location). This simplifying assumptions may not always capture reality. How to capture the inhomogeneous feature of earthquake in space is a challenge but a foundation problem, to solve this problem, this paper introduces a Body-Fitted Coordinates based method to represent the covariance structure of seismograph array. This method need not to set a reference point, and can be taken as an alternative way to reflect the variation of seismic ground motion in inhomogeneous field.

Just as discussed above, the coordinate (x_{1i}, x_{2i}) in composition map (C plane) can reflect the difference or correlation of seismic ground motion, here we establish the bivariate mapping function of the C plane (x_1, x_2) into the G plane (y_1, y_2) by the sampling point of station, and set a rectangle grid in the C plane. By the mapping function we mapping the grid in the C plane into the G plane, the covariance structure of ground motion can be visualized through the density of grid in the G plane.

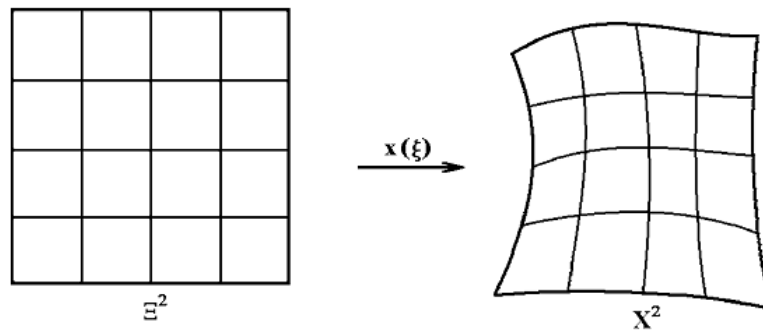


Figure 1. The mapping from the C plane (computational domain in left) to G plane (physical domain in right)

The main outline of this method can be described ad follow:

- (1) Calculate the DTW of the position (y_{1i}, y_{2i}) (G plane)
- (2) Calculate the Composition map by the MDS we get the coordinates (x_{1i}, x_{2i}) in the C plane ($i = 1, 2, \dots, N$).
- (3) Establish the relationship of mapping by a bivariate function f :

$$\begin{pmatrix} y_{1i} \\ y_{2i} \end{pmatrix} = f \begin{pmatrix} x_{1i} \\ x_{2i} \end{pmatrix} \quad (9)$$

We define a roughness criterion:

$$J_i(f) = \int_{R^2} \left[\left(\frac{\partial^2 f_i}{\partial x_1^2} \right)^2 + 2 \left(\frac{\partial^2 f_i}{\partial x_1 \partial x_2} \right)^2 + \left(\frac{\partial^2 f_i}{\partial x_2^2} \right)^2 \right] dx_1 dx_2 \quad (10)$$

and compute the bivariate function $f = (f_1, f_2)$ to minimize

$$L = \sum_{j=1}^2 \sum_{i=1}^N (y_{ij} - f_j(x_{ij}))^2 + \lambda (J_1(f_1) + J_2(f_2)) \quad (11)$$

for specified smoothing parameter λ . Where R^2 is the domain of interest.

(4) Define the point on the boundary line of the rectangle grid in the composition map. And use the Mapping Function f to get the correspond point in the G space: by the interpolating spline we can generate curve line and establish the grid.

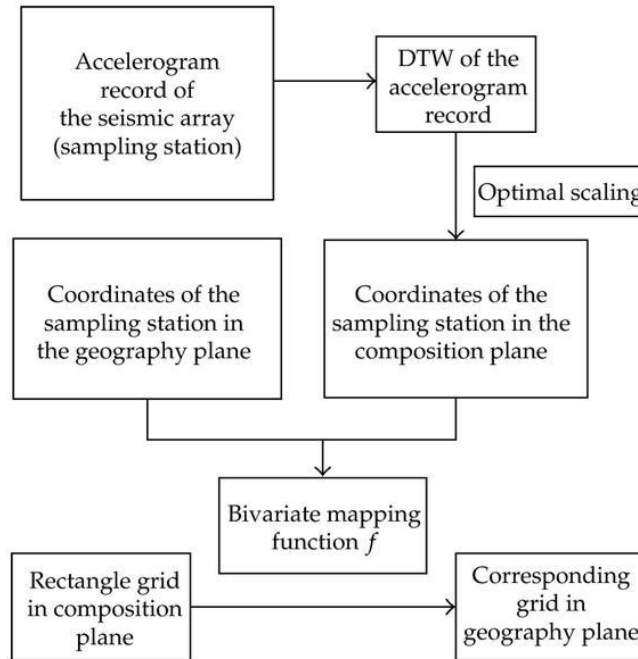


Figure 2. outline of the method

3. RESULTS AND DISCUSSIONS

In this section we will do a preliminary analysis of seismograph array data by using the non-parametric estimation method proposed. This example manifests a somewhat extreme, but easily explained, form of nonstationary in the spatial covariance structure of the earthquake. The earthquake ground motion data is from the event-40 recorded by Smart-1 array which is located in Lotung, Taiwan. The stations selected are located in a two-dimensional surface array consisting of a center C00 and three concentric circles (inner I, middle M, and outer O) with radii of 200, 1000, and 2000m, respectively, each with 12 strong-motion seismographs having common time base, (see Fig 3). The station applied in this paper is I6, I9, I11, M10, M07, M03, O05, O06, O10, M06. The epicentral direction is 15° to the north. The magnitude (M^L) of the earthquake is 6.5, and the location of earthquake source relative to the center station C00 is 22 km, the depth of the source is 10 km. The recorded motion has a sampling rate of 100Hz and a total of 1201 values. The computation uses an Ormsby filter to eliminate long periods from the acceleration recordings. The cut-off and roll-off frequencies are $f_c = 0.167\text{Hz}$ and $f_T = f_c + 0.03 = 0.197\text{Hz}$.

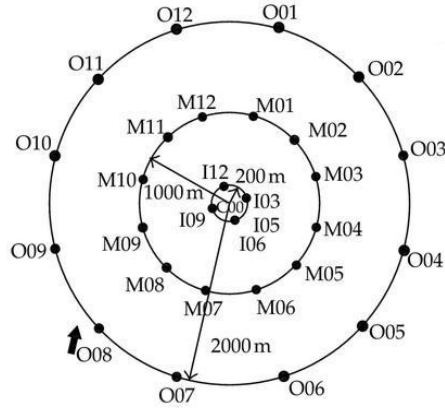


Figure 3. The array of smart-1

By the aforementioned method in 2.1.1 we compute the observed proximity matrix $[\delta_{ij}] = DTW(i, j)$ of sampling, in table 1.

Table 1. The DTW of sampling station

	I6	I9	I11	M10	M07	M03	O05	O06	O10	O11	M06
I6	0	0.8036	0.8642	1.0437	1.0759	1.1891	1.1258	1.0009	1.2965	1.3018	1.1234
I9	0.8036	0	0.5779	0.939	1.197	0.9975	1.0781	1.0136	1.1992	1.333	0.9159
I11	0.8642	0.5779	0	0.7755	1.2855	1.1439	1.0705	1.0546	1.3785	1.3817	1.0058
M10	1.0437	0.939	0.7755	0	1.2088	1.2764	0.9947	1.0047	1.408	1.5089	1.1313
M07	1.0759	1.197	1.2855	1.2088	0	1.226	0.9112	1.0594	1.1046	1.5046	1.0792
M03	1.1891	0.9975	1.1439	1.2764	1.226	0	1.2636	1.3458	0.8758	1.3413	1.1292
O05	1.1258	1.0781	1.0705	0.9947	0.9112	1.2636	0	1.4143	1.4136	1.4847	1.414
O06	1.0009	1.0136	1.0546	1.0047	1.0594	1.3458	1.4143	0	1.2622	1.4257	1.0249
O10	1.2965	1.1992	1.3785	1.408	1.1046	0.8758	1.4136	1.2622	0	1.191	1.2415
O11	1.3018	1.333	1.3817	1.5089	1.5046	1.3413	1.4847	1.4257	1.191	0	1.3824
M06	1.1234	0.9159	1.0058	1.1313	1.0792	1.1292	1.414	1.0249	1.2415	1.3824	0

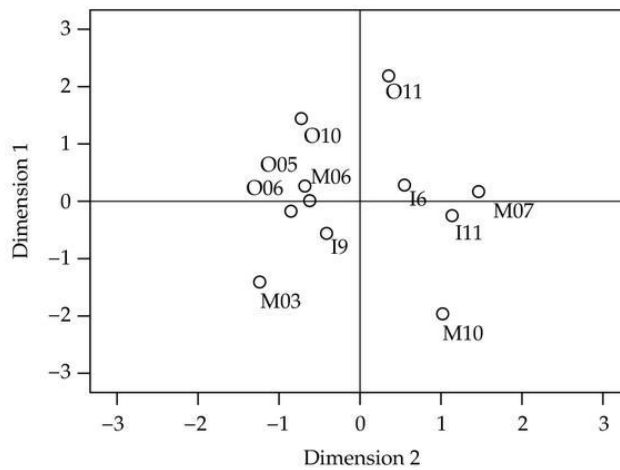


Figure 4. Coordinates of sampling station in the C plane

By the aforementioned method in section 2.1.2 and the proximity matrix $[\delta_{ij}]$ in Table2, we can get the

coordinates of the sampling station in the C plane, in Fig 4. For example stations record that are located close to each other in C plane are perceived as being similar such as I06 and I11. In contrast, stations positioned far away from each other, indicate a large difference in perception, such as I06 and M10.

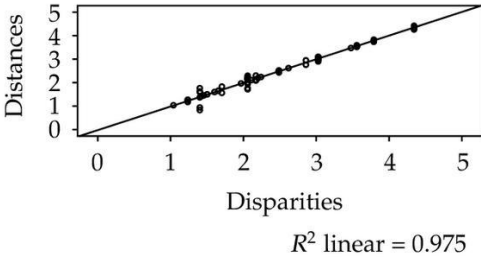


Figure 5. Disparities versus MDS interstation Distances for the seismograph array date

Figure 5 is the distance-disparities scates map, it can be observed that the scatter plot of distances d_{ij} in the C plane versus the observed spatial dispersions is nearly a straight line, It shows that this two-dimensional MDS solution accurately reflects the date in the matrix $\Delta = [\delta_{ij}]$, and the optimal metric have better linear relation than the Classic metric method.

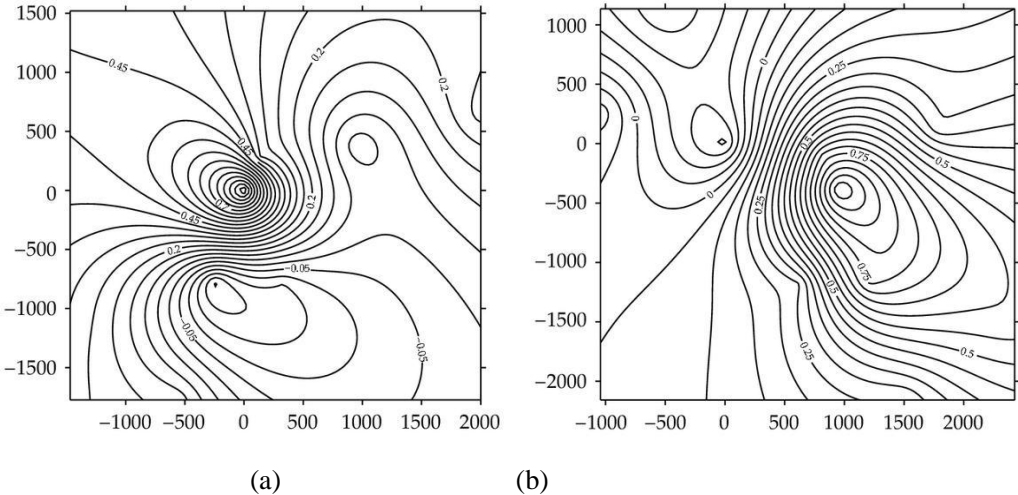


Figure 6. Contour plot of correlation coefficient of sample record for event 40:
 (a) take the I09 as the reference point (b) take M05 as the reference point (the unit of the coordinates is m).

Estimated Contours plots of the correlation coefficient are presented based on the strong motion recorded by smart-1 array in Fig 6. What information conveys in Fig6 is that even in the alluvium alloy where seismic array located, the covariance structure of the earthquake is inhomogeneous in space, or varies from location to location. The different choose of reference point may get different result, so the contours plot of the correlation coefficient is not a suitable way of represent the spatial variation of seismic ground motion.

Fig 7 is the figure by the method in this paper; it depicts the thin-plate spline mapping between the G plane and C plane representations of monitoring stations using the image of a rectangular of points located on C plane. Different from Contour plot in Fig 6, it needs not to choose the reference point. The clearest feature in Fig 7 is that there is a relatively denser spacing of curves in the north-east, it is the epicentral direction. The area where the grid (spacing of curves) concentrated is the place where the correlative structure of earthquake changes significantly, the denser of grid means greater variation, and vice versa. It should be emphasized that this paper concentrates to illustrate how to apply our method to despite the corrective structure of earthquake ground motion, a more exact conclusion need

more sufficiently dense spatial sampling.

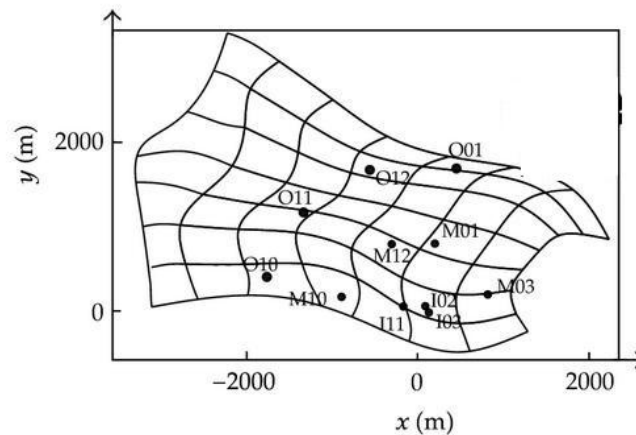


Figure 7. The Grid for the spatial correlation structure of earthquake ground motion in geography space.

4. CONCLUSIONS AND REMARKS

This study addresses the topic: The description and estimation of spatial variation of seismic ground motions. A Non-parametric method to the estimation and graphical depiction of the local spatial correlation for the earthquake ground motion is presented. The method paves a way how to use the multi-dimension scaling and Boy-fit grid to represent the variation of ground motion. It can be a promising tool in detecting the correlation grid structure of spatial ground motion; further research includes the investigation of the influence, on the spatial correlation structure, of the seismologic parameters such as the epicenter the magnitude and location.

Such covariance structure estimate by this method can be used to reveal the inhomogeneous feature of earthquake, and help design large-span structure, for example a suggested way to do so is: modify the traditional correlation model. By observing the density of the grid in the geography space we can introduce different weight coefficients at different locations. In this way we can produce more properly correlated motions in space for the design of engineering structure with character of spatial extension. The problem how to establish the relationship between the grid by our method and weight coefficients of correlation model is a question should be addressed in future work before the application of our methods to problems of artificial earthquake simulation.

The covariance estimates can also be used for some monitoring network design problems. For example: grid concentrated is the place more monitoring points (stations) should be placed, and vice verse.

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