



EARTHQUAKE GROUND SHAKING HAZARD MAPS FOR THE STATE OF MONTANA

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SUMMARY

Seismic hazard maps are a useful tool for regulatory agencies that are responsible for the safety of important and critical facilities particularly if their area of responsibility is large and geologically and seismically diverse and if they manage a large inventory. Western Montana, located in the northern Intermountain region of the U.S., is characterized by abundant late-Quaternary Basin and Range normal faulting and historical seismicity. In contrast, eastern Montana is similar to other portions of the Great Plains in the central U.S., with only two known faults of possible late-Quaternary age and a relatively low level of seismicity.

Because of the potential earthquake threat to dams in the state, we have developed a set of probabilistic earthquake ground motion maps for the Montana Department of Natural Resources and Conservation Dam Safety Program. The 18 statewide maps display peak horizontal acceleration and 0.2 and 1.0 sec horizontal spectral acceleration for approximate return periods of 500, 2500, and 5000 years (exceedance probabilities of 10%, 2%, and 1% in 50 years, respectively). The maps display ground motions for two site conditions: soft rock and the ground surface. Ground motions calculated at the ground surface include site response effects for the areas underlain by unconsolidated sediments, such as the major late Cenozoic basins in western Montana.

Ground motion attenuation also differs significantly in the western and eastern parts of the state reflecting their contrast in geologic and tectonic characteristics. The lack of strong motion data for Montana and surrounding regions required the use of stochastic numerical modeling to develop region-specific attenuation relationships. These relationships were used together with published attenuation relationships

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to estimate the map ground motions. Amplification factors were then used to modify the rock motions and hence incorporate site response into the hazard maps. These factors were based on five generalized geologic site response categories (hard rock, soft rock, Quaternary alluvium, Quaternary lacustrine deposits, and Quaternary glacial till) and were adopted from similar Salt Lake Valley and California-based categories because insufficient subsurface geologic and geotechnical data are available for Montana. Depth-dependent site amplification factors were used for the basins in western Montana.

The resulting hazard maps for an uniform site condition of soft rock show the highest hazard is concentrated along the most active faults. For example, at a return period of 5000 years, the highest peak horizontal accelerations reach upwards of 1 g in the vicinity of the active Centennial fault along the Montana-Idaho border. Other relatively moderate hazard areas occur in the vicinities of the Mission, Canyon Ferry, Madison, Emigrant, and Red Rock faults. Areas away from the more active faults in western Montana are characterized by values of 0.10 to 0.30 g. Eastern Montana is characterized by peak horizontal accelerations of less than 0.10 g at a return period of 5000 years due to the absence of active faults and a low-level of seismicity.

INTRODUCTION

Western Montana is characterized by abundant late-Quaternary Basin and Range normal faulting and historical seismicity (Figure 1). It includes the Intermountain Seismic Belt (ISB), a zone of elevated seismicity, and the seismically and volcanically-active Yellowstone region [1]. Paralleling the southwestern border of the state is the Centennial Tectonic Belt (CTB), also a zone of significant seismicity [2]. The largest historical event was the 1959 moment magnitude (**M**) 7.3 Hebgen Lake earthquake located just west of Yellowstone National Park. In contrast, eastern Montana, similar to other portions of the Great Plains in the central U.S., has only two known faults of possible late-Quaternary age. Seismicity is also at a relatively low level although the largest event has been about **M** 5.5 (Figure 1).

Because of the potential earthquake threat to dams in the state, we have developed a set of probabilistic earthquake ground motion maps for the Montana Department of Natural Resources and Conservation (DNRC) Dam Safety Program. Sponsored by the National Dam Safety Act, the statewide maps display peak horizontal acceleration and 0.2 and 1.0 sec spectral acceleration for 10%, 2%, and 1% in 50 years exceedance probabilities (return periods of 500, 2500, and 5000 years, respectively). This range of exceedance probabilities was selected to cover the range of hazard categories of Montana dams based on downstream risk. Ground motions are calculated for a rock site condition and at the ground surface; the latter includes site response effects for the areas underlain by unconsolidated sediments. A total of 18 maps were developed for the three exceedance probabilities, three spectral accelerations, and two site conditions. The maps are being used by DNRC (1) to check site-specific seismic hazard evaluations required of all owners of high hazard dams, (2) as the basis for hazard assessments for all low hazard dams, and (3) to prioritize the state's dams in terms of hazards and risk, for potential remediation. A complete description of the methodology, inputs and results of the hazard analysis, and the maps can be found in Wong *et al.* [3].

In the development of hazard maps, the effects on ground motions from source, path, and site conditions need to be incorporated into seismic hazard analysis. In particular, near-surface site amplification is expected to be a significant factor affecting the ground shaking hazard at valley (basin) sites in western Montana. Observations of near-field ground motions also suggest that rupture directivity and hanging wall effects may also be significant in areas adjacent to active faults in the state. The estimates of strong ground

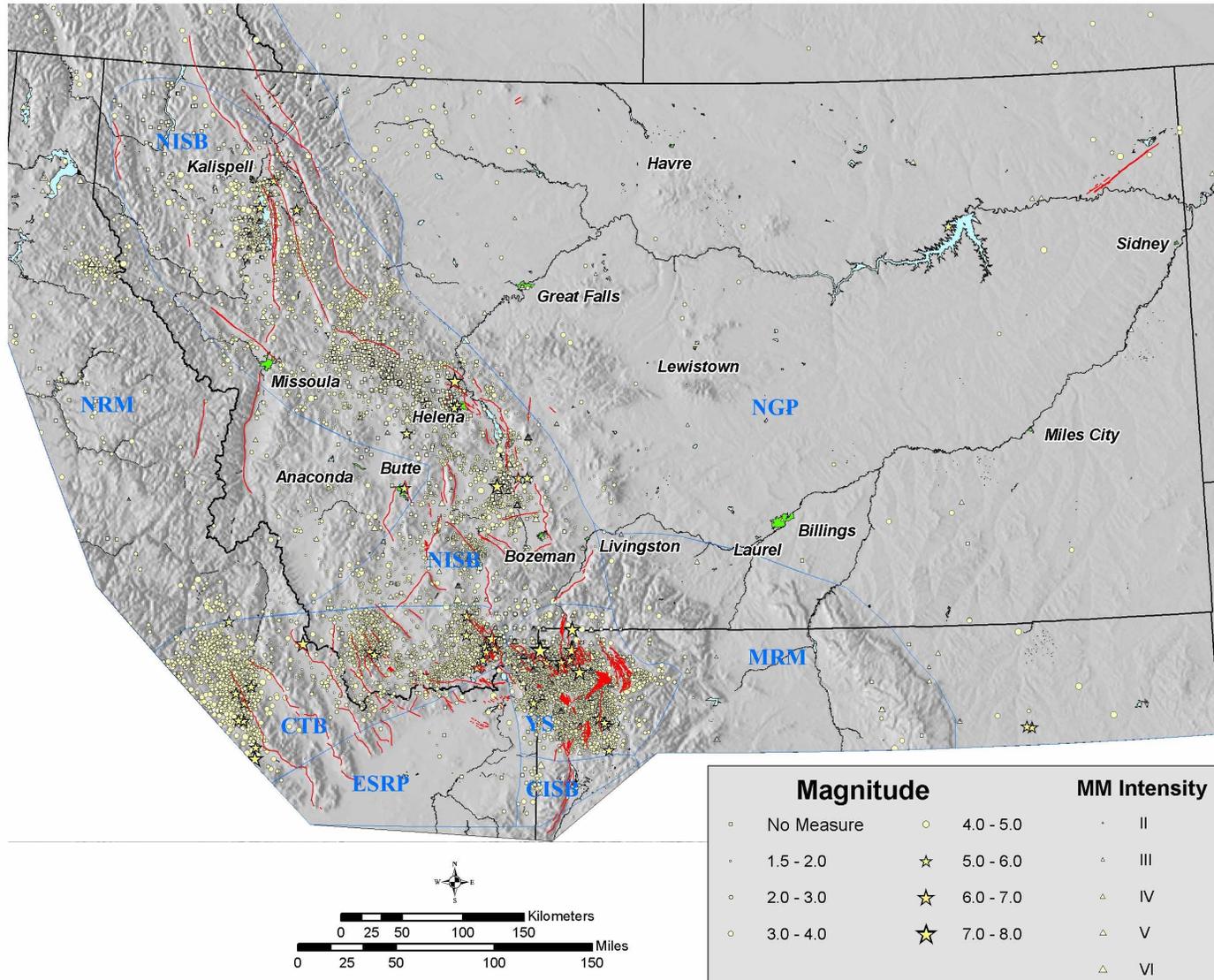


Figure 1. Historical seismicity (1809 to 2001), Quaternary faults, and regional seismic source zones in the State of Montana

shaking for Montana displayed on the maps and described in this paper are based on the most up-to-date albeit limited information on seismic sources, crustal attenuation, and near-surface geology. Because of this limited data, there are large uncertainties associated with earthquake ground motion prediction in Montana. These uncertainties must be considered in the use of our maps.

The hazard maps are intended to illustrate the intensity and variability of ground shaking within the state for the three exceedance probabilities. In addition to DNRC's use of the maps, we hope the maps will increase general public awareness of earthquake hazards, and will aid in regional planning, emergency preparedness, and loss estimation. Although we believe the maps represent the state-of-the-art in ground motion modeling, the maps are not intended to be used directly in engineering design. Various codes (e.g., International Building Code) define minimum design ground motion levels for buildings, bridges, and other structures, and commonly reference other maps to determine these design levels. Our maps do not replace these design maps, but can be used to compare with code-based design, and to evaluate the need for increasing design levels if indicated.

SEISMOTECTONIC SETTING AND HISTORICAL SEISMICITY

Western Montana is one of the most seismically active regions within the interior of the western U.S. Seismicity in western Montana is largely concentrated within the northern ISB, the CTB, and the Yellowstone region [1] (Figure 1). In contrast, eastern Montana has a much lower, albeit not insignificant, level of seismicity due to its location within the comparatively tectonically stable Northern Great Plains Province.

The ISB extends from northwestern Arizona through central Utah, along the Idaho-Wyoming border and terminates in northwestern Montana [1]. The deformational processes occurring along the ISB are principally in response to ongoing tectonic extension as observed throughout the Basin and Range Province. In general, the ISB is characterized by late-Quaternary normal faulting, diffuse shallow seismicity generally occurring in the top 15 km of the crust and episodic scarp-forming earthquakes (M 6.5 to 7.5) [1]. Within the northern ISB in western Montana, seismicity is diffuse with concentrations of small magnitude ($M \leq 4.0$) events corresponding to aftershocks of $M \geq 6.0$ events (Figure 1).

Stickney and Bartholomew [2] define the southern margin of the northern Basin and Range Province as the CTB (Figure 1). The CTB represents the northern arm of the parabolic distribution of seismicity and late Quaternary faulting centered on Yellowstone and flanking the Snake River Plains [4]. The CTB, which is part of the northern ISB, is the most tectonically and seismically active portion of the northern Basin and Range Province and was the site of the 1983 M 6.8 Borah Peak, Idaho earthquake. Stickney and Bartholomew [2] suggest that earthquakes are more frequent and larger in the CTB than in the northern ISB.

Within western Montana, there have been at least 18 earthquakes of $M \geq 5.0$ from about 1809 to the present. The largest of these earthquakes, the 1959 Hebgen Lake event is the only earthquake to have ruptured the ground surface. The 1925 M 6.6 earthquake has been located in the southern end of Clarkston Valley, although the responsible fault remains unknown to this day. Other significant earthquakes in western Montana include events in Beaverhead County in 1897 (Richter magnitude M_L 6.4), near Helena in 1935 ($M \sim 6\frac{1}{4}$), Virginia City in 1947 (M_L 6.3), and Yellowstone in 1975 (M_L 6.4).

The most significant historical earthquake known in eastern Montana was an event of estimated M 5.5 that occurred near the border between northeastern Montana and Saskatchewan, Canada. Bakun and

Stickney [5] suggest that the source of the earthquake could be the Brockton-Froid fault, one of the two postulated late-Quaternary faults in eastern Montana (Figure 1).

METHODOLOGY AND INPUT TO HAZARD CALCULATIONS

There were five principal tasks in this study: (1) seismic source characterization; (2) definition and characterization of geologic site response categories and assignment of amplification factors; (3) seismic attenuation characterization; (4) probabilistic ground motion calculations; and (5) map development.

Seismic Source Characterization

The first step in any assessment of earthquake ground-shaking hazards is a characterization of the seismic sources that will produce ground motions of engineering significance at the site or area of interest. Seismic source characterization is concerned with three fundamental elements: (1) the identification, location, and geometry of significant sources of earthquakes; (2) the maximum size distribution of earthquakes for each source; and (3) the rate at which different size earthquakes occur in each source. Parameters needed are fault location, geometry, orientation, sense of slip, maximum earthquake magnitude (M_{\max}), and recurrence model and rate. Uncertainties in the seismic source parameters were incorporated into the probabilistic seismic hazard analysis using a logic tree approach. In this procedure, values of the source parameters are represented by the branches of logic trees with weights that define the distribution of values. In general, three values for each parameter were weighted and used in the analysis.

In a probabilistic hazard assessment of earthquake ground motions, all seismic sources that can generate significant ground shaking at a site (generally those within a distance of 100 to 200 km in the western U.S.) are characterized. Two general types of seismic sources were considered in the probabilistic hazard analysis: active or seismogenic faults and an areal source zone.

Quaternary Fault Sources

Our seismic source characterization model for this analysis included 92 potential fault sources (Figure 1). We included all known faults longer than 5 km with evidence for repeated Quaternary movement within Montana and extending out to 50 km beyond the border. We also considered more significant faults out to 100 km, including longer more active faults (e.g., Teton, Lemhi, and Lost River faults). We used a variety of data sources but relied heavily on Quaternary fault compilations by Haller *et al.* [6], Stickney *et al.* [7], and Machette *et al.* [8]. Several previous seismic hazard analyses performed by us in the region were also used (e.g., Woodward-Clyde Federal Services *et al.* [9], Wong *et al.* [10, 11] as well as fault data from recent U.S. Geological Survey, U.S. Bureau of Reclamation, Montana Bureau of Mines and Geology (MBMG), and Montana State University studies (e.g., Pierce *et al.* [12], Hecker *et al.* [13], Haller *et al.* [14], Anderson and LaForge [15, 16], Ruleman [17], Liebli and Lageson [18]). We also contacted various individuals about ongoing and unpublished studies (e.g., S.M. Vuke, MBMG, written communication, 2003; J.E. Adams, Geological Survey of Canada, written communication, 2002; Breckenridge *et al.* [19]). Unfortunately, most of the Quaternary faults in Montana (> 85%) have not been studied in any detail and very few data are available to develop rupture models and constrain slip rates. Our fault parameters are fully tabulated and documented with references, which are not included here, but will be included in Wong *et al.* [3] to facilitate updating the model as new data becomes available.

Where the data permit, we have attempted to consider and accommodate the structural variations that are potentially significant to the hazard analysis by including a variety of rupture behavior models and fault geometries in our source characterization. All faults are dominantly normal-slip faults that we modeled as single planes, zones (multiple planes), or curvilinear surfaces. Most faults are included as single independent (unsegmented) planar sources, unless the available data suggest otherwise. Zones of faults are modeled as multiple, distributed, parallel fault planes that do not coseismically rupture as in the linked

model, but rather have the assigned moment rate distributed evenly among the planes. Most of zones of faults are associated with volcanism and generally have low rates of activity (e.g., Elephant Back fault zone, Mirror Plateau faults, and Mallard Lake resurgent dome faults).

Alternative rupture behavior models to the single-plane, independent fault include linked faults, segmented faults, and zones of faults. Potentially linked faults may experience coseismic rupture (either along or across strike), whereas portions of potentially segmented faults may rupture independently of each other. Examples of potentially linked faults include the newly discovered Toston fault [20] and the Canyon Ferry fault [21], the several faults of the 1959 Hebgen Lake earthquake rupture, and the Georgia Gulch and Tobacco Root faults. Some faults show compelling evidence for being segmented, where relatively persistent segment boundaries have apparently confined prehistoric surface ruptures to particular sections of the faults (e.g., Lost River, Mission, and Centennial faults). For other faults, the evidence is more ambiguous as to whether persistent rupture segment boundaries exist (e.g., Blacktail, Emigrant, and Teton faults). Finally, we note that the rupture behavior of most of the faults in the state is poorly understood and may actually be more complex than our simplifying assumptions, but we have attempted to address uncertainties that are particularly significant to the hazard in Montana, given the available data.

The youngest most active faults and fault segments (i.e., late Quaternary faults with slip rates ≥ 0.1 mm/yr) were modeled as curvilinear surfaces. To model the curvilinear nature of normal faults, we digitized the primary, most-active fault trace and projected these curves down-dip using a weighted mean strike. Thus, these simple curvilinear surfaces retain a constant dip and do not accommodate complexities like listric faults (i.e., decreasing dip with depth). All surfaces, planar and curvilinear, extend the full depth of the seismogenic crust, and so fault dips are averages estimated over the full depth of the seismogenic crust. All faults as well as regional source zones (see “Background Seismicity”) were assigned a maximum seismogenic depth of 13, 15, and 17 km, weighted 0.2, 0.6, and 0.2, respectively. For most typical range-bounding normal faults, preferred dips are assumed to be 55° unless fault-specific data indicate otherwise.

Preferred M_{\max} ranged from M 6.2 to 7.3. These values were estimated using the empirical relationships of Wells and Coppersmith [22] for all fault types. Magnitudes were generally solely based on fault lengths, but we also considered displacements per event where data were available. In our analysis, we include distributions of ± 0.3 around the preferred M_{\max} to account for the various epistemic uncertainties in determining M_{\max} for faults (e.g., dePolo and Slemmons [23]). These include uncertainties associated with the regression relations used and the input parameters to those relations, insofar as uncertainties in maximum rupture lengths and/or displacements per event were not explicitly included.

Probabilities of activity ranged from 0.2 to 1.0. In assigning probabilities of activity for each fault source, we considered both the likelihood that it is structurally capable of independently generating earthquakes (seismogenic), and the likelihood that it is still active within the modern stress field. We incorporated many factors in assessing these likelihoods, such as: orientation in the modern stress field, fault geometry (length, continuity, depth extent, and dip), relation to other faults, age of youngest movement, rates of activity, geomorphic expression, amount of cumulative offset, and any evidence for a non-tectonic origin. Faults with definitive evidence for repeated Quaternary activity were generally assigned probabilities of being active and seismogenic of 1.0. Exceptions include faults that may be secondary and dependent on other faults, faults or suspected fault features that may have a non-seismogenic origin, and faults that may be too short to independently generate large earthquakes (≤ 10 km length). The probability of activity for faults that do not show definitive evidence for repeated Quaternary activity was individually judged based on available data and the criteria explained above.

We considered truncated-exponential, characteristic, and maximum-magnitude recurrence models, with weights dependent on the fault length, type of data used to calculate activity rates, and type of rupture model. The truncated exponential recurrence model is the traditional Gutenberg-Richter exponential frequency-magnitude relationship that is truncated at the maximum magnitude. As defined by Schwartz and Coppersmith [24], the characteristic recurrence model was based on paleoseismic observations of similar-size displacements per event on faults, implying that surface faulting events more typically exhibited a “characteristic” magnitude rather than a full range of magnitudes, including more frequent moderate-sized (M 6 to 7) events, as inferred from extrapolating the historical record of seismicity and using the exponential model. Thus, the characteristic model predicts fewer moderate-size events and generally results in lower hazard than the truncated exponential model. We use the characteristic model of Youngs and Coppersmith [25]. The maximum magnitude model is an extreme version of the characteristic model and assumes that a fault (or fault segment) only ruptures in its entirety in characteristic-sized events, and smaller events do not occur (e.g., Wesnousky [26]).

Observations of historical seismicity and paleoseismic investigations worldwide suggest that characteristic behavior is more likely for individual faults, whereas seismicity in zones best fits a truncated exponential model [24]. Therefore, except for zones of faults, we generally favored the characteristic model of Youngs and Coppersmith [25] for all fault sources by giving it weights of 0.6 to 0.8. We assigned equal weights to the exponential and characteristic models for zones of faults. We assigned a slightly higher weight to the maximum-magnitude model than to the exponential model for longer, segmented faults (0.2 versus 0.1), but for shorter independent faults we assigned both models an equal weight of 0.2. This choice was based on the idea that as faults develop, become longer and eventually segmented, their behavior may evolve to become less exponential and more characteristic (e.g., Wesnousky [27, 28], Stirling *et al.* [29]).

We incorporated all available intermediate- (≤ 1.6 Ma) and short-term (≤ 130 ka) slip rate data in characterizing rates of fault activity, but we generally preferred short-term data when available. Weighted mean slip rates ranged from 0.0006 mm/yr (Lake of the Woods fault) to 1.7 mm/yr (Teton fault). Although we only used slip rates here, we usually prefer using recurrence intervals instead of slip rates to characterize rates of activity [11]. However, recurrence intervals are generally lacking for Montana faults, and in our judgment, the quality of the available data are inferior to the available slip rate data. We converted vertical slip rates to net (fault parallel) slip rates for most faults by assuming 100% dip slip and the preferred fault dips for each individual fault. Additionally, wherever possible we attempted to calculate or adjust for along-strike average slip rates. Variations of displacements along strike can significantly affect the calculation of slip rates [30], but unfortunately very few faults have enough data to calculate average rates for the entire fault (e.g., Teton fault). More typically we found only a few data points for one or two sites along the fault (e.g., Canyon Ferry fault) or no fault-specific data at all (e.g., Boulder River Valley western border fault). In the latter case, we assumed slip rate distributions to be the same as a similar nearby structure, taking into account such factors as style of deformation, geomorphic expression, and age of youngest movement.

Background Seismicity

The hazard from background (floating or random) earthquakes that are not associated with known or mapped faults needs to be incorporated into the probabilistic hazard analysis. Earthquake recurrence estimates in the study region and M_{\max} are required to assess the hazard from background earthquakes. In most of the western U.S., particularly the Basin and Range Province, the M_{\max} for background earthquakes usually ranges from M 6 to $6\frac{1}{2}$. Repeated events larger than these magnitudes probably produce recognizable fault- or fold-related features at the earth’s surface (e.g., Doser [31], dePolo [32]). In this

study, we adopt values of $M 6$ to $6\frac{1}{2} \pm \frac{1}{4}$ depending on the seismotectonic setting, i.e., regional source zone. Recurrence estimates were calculated for the nine regional source zones considered in this study: the northern and central ISB (NISB and CISB), CTB, Yellowstone region (YS), Northern and Middle Rocky Mountains (NRM and MRM), Northern Great Plains (NGP), and the eastern Snake River Plain (ESRP) (Figure 1).

In addition to the traditional approach of using areal source zones (assuming uniformly distributed seismicity), Gaussian smoothing [33] was used to address the hazard from background earthquakes in the probabilistic analysis. In this approach, we smoothed the historical background seismicity to incorporate a degree of stationarity, using a spatial window of 15 km.

Geologic Site Response Units and Amplification Factors

In order to quantify the site response of soil and unconsolidated sediments, a shear-wave velocity profile, depth to a reference rock datum, and dynamic degradation curves (both shear modulus reduction and damping) are required to define geologic site response categories. Based on these site response categories, frequency- and strain-dependent amplification factors can be computed as a function of input ground motions. Unfortunately, there were insufficient geologic and geotechnical data to characterize the near-surface geology in the Montana in sufficient detail to calculate region-specific amplification factors. Shear-wave velocity data in the state were sparse with only a few known measurement sites. Depths to reference rock were only known for the major basins in western Montana. Thus in the absence of region-specific data, we defined five generic geologic site response categories: hard rock, soft rock, Quaternary alluvium, Quaternary lacustrine deposits, and Quaternary glacial till. Based on a surficial geologic map of the state that we assembled from maps from the MBMG and USGS, lithologic units were assigned to one of the five site response categories. Differences in lithologic unit classification and discontinuities of map units across quadrangle boundaries were a major source of uncertainty on the resulting state map.

Since region-specific amplification factors could not be computed due to a lack of data, we adopted factors based on comparisons with the surficial geology in the Salt Lake Valley [34] and California [35]. These factors were developed using an equivalent-linear site response approach coupled with the stochastic numerical modeling technique [36]. To account for uncertainties in this approach, the amplification factors for broadly similar geologic categories were combined and enveloped. With the exception of 18 major Quaternary basins in western Montana, the thickness of unconsolidated sediments in the state is generally less than 30 m. Within the basins, the sediments are up to 300 m thick although the uncertainties in thickness estimates are large. Depth-dependent amplification factors for the basins were calculated by enveloping factors over ranges of thicknesses (e.g., 50 to 100 m and 100 to 200 m). Note because of the division of rock into soft and hard, amplification factors were adopted that allowed modification of ground motions from soft rock into hard rock and vice versa.

Ground Motion Attenuation Characterization

To characterize the attenuation of ground motions, we used attenuation relationships appropriate for soft rock sites in the western U.S. and hard rock sites in the Midcontinent, and a stochastic numerical modeling technique [36]. An important consideration in the selection of attenuation relationships is that western Montana is located in the extensional Basin and Range Province where normal faulting dominates and eastern Montana lies within the compressional Midcontinent. It has been increasingly recognized that earthquakes in extensional tectonic regimes produce lower ground motions than events in compressional/strike-slip regimes for the same magnitude and distance [37, 30, 38].

We used the following empirical soft rock relationships in this study for western Montana: Abrahamson and Silva [39], with normal faulting factors; Spudich *et al.* [40], which was developed from an extensional

earthquake strong motion database; Sadigh *et al.* [41]; and Campbell [42]. The latter two relationships are based primarily on California strong motion data and were included to more fully address uncertainty. None of these relationships are specific to Montana or the Basin and Range Province due to the paucity of strong motion records. The relationships were weighted 0.40, 0.40, 0.10, and 0.10, respectively, based on our subjective judgment of the applicability of each relationship.

For eastern Montana, we used the central and eastern U.S. hard rock relationships of Atkinson and Boore [43], Somerville *et al.* [44], Toro *et al.* [45] for the Midcontinent, Campbell [46] and the single- and double-corner models of Silva *et al.* [47]. These relationships were weighted 0.12, 0.08, 0.12, 0.08, 0.40, and 0.20, respectively. It should be noted that due to the almost complete absence of strong motion records in the central and eastern U.S., these relationships have been developed based on numerical ground motion modeling.

To compensate for the lack of region-specific attenuation relationships, the stochastic ground motion modeling approach was used to develop such relationships [48] for both western and eastern Montana. The point-source version of the stochastic methodology [36] was used to model earthquakes of M 5.5, 6.5, and 7.5 in the distance range of 1 to 400 km. Uncertainties in stress drop, magnitude-dependent focal depths, the crustal attenuation parameters Q_0 and η , the near-surface attenuation parameter (κ), and the rock profile atop the crustal model were included in the computations of the attenuation relationships through parametric variations [48]. Ranges of magnitude-dependent stress drops appropriate for extensional and compressional regimes were used [49]. The P-wave crustal velocity model used for locating earthquakes in western Montana was adopted in the calculations. Inserted on top of this model is a generic western U.S. soft rock V_s profile developed from the database compiled by Pacific Engineering & Analysis. For eastern Montana, a crustal model based on the seismic investigations by Warren *et al.* [50] and McCamy and Meyer [51] was used to develop the stochastic attenuation relationships.

The uncertainty in ground motion attenuation was included in the probabilistic analysis by using the log-normal distribution about the median values as defined by the standard error associated with each attenuation relationship. Three standard deviations about the median value were included in the analysis.

GROUND MOTION CALCULATIONS AND MAP DEVELOPMENT

Ground motions were estimated for the probabilistic hazard maps using the previously described seismic source model and the attenuation relationships. Peak horizontal acceleration and spectral accelerations at spectral periods of 0.2 and 1.0 sec were displayed in map form using GIS.

Probabilistic Ground Motions

To calculate the probabilistic ground motions, we performed a comprehensive Cornell [52] hazard analysis using logic trees, employing the computer code HAZ32 written by Norm Abrahamson. All known seismic sources that could generate strong ground shaking in the state, were incorporated into the probabilistic analysis. Published and stochastic attenuation relationships, weighted 0.40 and 0.60, respectively, were used in the analysis to calculate the ground motion values. The mean probabilistic hazard was calculated at return periods of 500, 2500, and 5000 years.

Map Development

The ground shaking maps were produced using a vector- and raster-based GIS. Probabilistic ground motions on rock were calculated for the map area using a grid of points at variable spacings of 5 to 20 km. Each grid point was assigned to a site response category. Surface ground motions were calculated by multiplying the rock ground motions by the appropriate amplification factors. The amplification factors

for each grid point were selected based on the site response unit, depth to rock, and the input rock peak acceleration as described above. For a grid point on soft rock in western Montana, no amplification factor was applied since the ground motions were calculated for a soft rock condition. For a grid point on soft rock in eastern Montana, an amplification factor was applied to convert from the hard rock input motion to soft rock. For each map, the peak or spectral acceleration values were color contoured by interpolation in intervals of 0.10 g. The intent was to avoid implying a greater level of resolution and/or accuracy than was possible given the limitations of available geologic data.

MAPS AND RESULTS

The resulting hazard maps for a uniform site condition of soft rock show the highest hazard concentrated along the most active faults (Figure 2). At a return period of 5000 years, the highest peak horizontal accelerations reach upwards of 1 g in the vicinity of the active Centennial fault along the Montana-Idaho border. Other areas of relatively moderate hazard occur in the vicinities of the Mission, Canyon Ferry, Madison, Emigrant, and Red Rock faults (Figure 2). Areas away from the more active faults in western Montana are characterized by peak accelerations of 0.10 to 0.30 g. Eastern Montana is characterized by peak horizontal accelerations of less than 0.10 g at a return period of 5000 years due to the absence of active faults and a low-level of seismicity (Figure 2). The peak acceleration hazard maps at return periods of 1000 and 2500 years show a similar pattern to the 5000-year return period maps. The peak horizontal acceleration value range up to 0.30 to 0.40 g and 0.70 to 0.80 g on the two maps, respectively. The hazard maps for 0.2 and 1.0 sec horizontal spectral acceleration display the same general patterns as the peak acceleration maps. The maps for the ground surface have been completed but are not discussed herein. They show a more complex pattern of ground motions due to the site response effects of Quaternary alluvium located in basins throughout western Montana and the Quaternary glacial till in the northernmost portion of the state.

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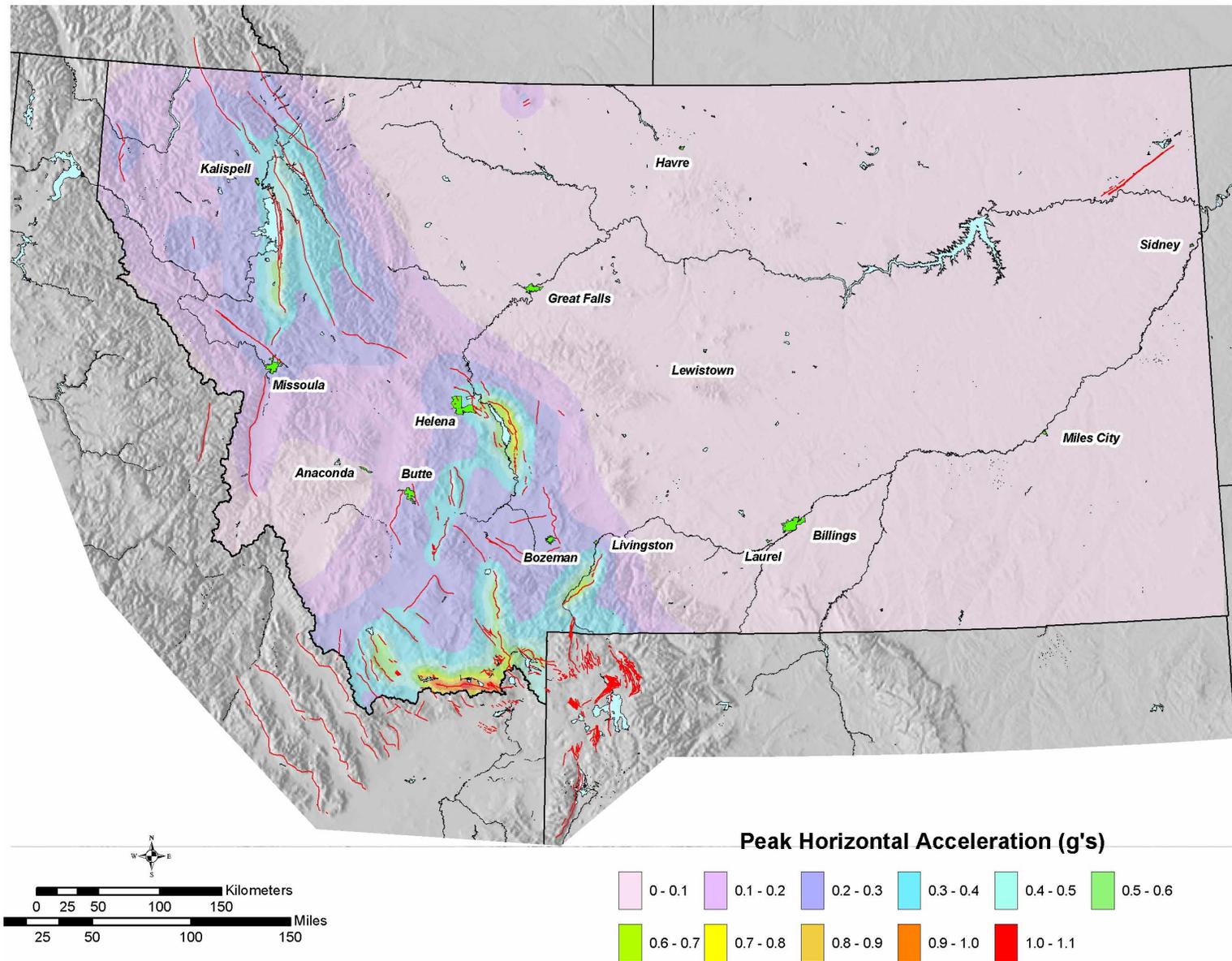


Figure 2. Peak horizontal acceleration (g) on rock for a 5000-year return period

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