

THE SEISMIC CHARACTERISTICS OF TURKISH EARTHQUAKES

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ABSTRACT

Turkey is a well-known earthquake-prone country and its earthquakes are generally of in-land type. The authors developed a data-base for Turkish Earthquakes for the last 2000 years and attempted to find out interrelations among seismic characteristics of these earthquakes. The data-base is named TURDEP and it is coded using a soft wear DBASE on a personal computer. Data files to study the interrelations among seismic characteristics are created using a program coded in BASIC from the data-base data. A soft-wear code called NGRAPH is used for graphical presentations of data. This integrated data-base system is described and seismic characteristics of Turkish earthquakes and their interrelations are presented and discussed.

KEYWORDS

Turkey; earthquakes; seismicity; characteristics; faults; data-base

INTRODUCTION

Turkey is a well-known earthquake-prone country and its seismic past is well documented (Ergin et al. 1967, Ayhan et al. 1981, Eyidoğan et al. 1990, Gençoğlu et al. 1990, Bayülke et al. 1990). Most of earthquakes are concentrated along the North Anatolian Fault (NAF), East Anatolian Fault (EAF), North East Anatolian Fault (NEAF) and West Anatolian Fault System (WAFS) as a result of north-ward motion of Arabian plate and African Continent and these earthquakes are generally of inland type which are more destructive than off-shore earthquakes even their magnitude could be smaller (Şengör 1980, Barka and Gülen 1987, Barka and Kadinsky-Cade 1988, Toksöz et al. 1979, Aydan and Paşamehmetoğlu 1994) (Figure 1).

The authors developed an integrated data-base system for Turkish Earthquakes for the last 2000 years with the purpose of finding out interrelations among their seismic characteristics. The data-base is named TURDEP and it is developed on a personal computer. In this article, this integrated data-base system is described and seismic characteristics of Turkish Earthquakes and interrelations among them are presented and discussed in order to provide information for seismic design of engineering structures as well as for forecasting earthquakes.

INTEGRATED DATA-BASE SYSTEM

Data-base System

The data base was named *TURDEP*, which stands for the first three letters of Turkish Earthquakes in Turkish, and was developed on a personal computer using a soft-wear *DBASE*. The items of the data-base system were chosen as follows:

- Group 1: Name, location (Latitude (Lat), Longitude (Lon)), Depth (Hd)
- Group 2: Magnitude $(M_b, M_s, M_{\rm ercalli} Intensity I_o)$
- Group 3: Peak ground acceleration (Acmax), dominant period(Tper), ratio of maximum vertical acceleration to maximum horizontal acceleration (Av/Ha)
- Group 4: Time (year (Year), month (Mo), day (Da), hour (Ho), minute (Mi))
- Group 5: Characteristics of rupture plane (dip direction (Ddr), dip (Dip), slip direction (Sdr), surface rupture length (Len), maximum relative displacement (Dis))
- Group 6: Type of rupturing (Strike-slip (Slst), Normal faulting (Norf), Reverse faulting (Revf))
- Group 7: Earthquake grouping (North Anatolian Fault (NAF), North East Anatolian Fault (NEAF), Central Anatolian Fault System (CAFS), East Anatolian Fault and Bitlis-Zagros Thrust Zone (EAFBZ), West Anatolian Fault System (WAFS))

Table 1 shows an example list of entries of the data-base system.

Table 1. A list of entries of data-base system

	Name of Earthquake	Location			Magnitude			Time Tper Year Mo Da Ho Mi Av/A							Rupture plane				Rupture Typ			Group	
No		Lat	Lon	Hd	Mb	Ms	lo	Acmax	Tper	Year	Mo	Da	Но	Mi	Av/Ah	Ddr	Dip	Sdr	Ler	Di	Typel	Type2	
101	KAS-PETHIYE	36 11	29.20	22 0	* A	6 9			0.00	1040		2.											• • • • •
102	CONEN	40.11	27 50	4.0	E 2	5 7	, 'T	0.0	0.00	1909			43	14	0.00	100	14	96	0		Revf		WAE
103	DEMIRCI .COPDES	30.10	20 45	12.0	3.0	9.1	- 4	0.0	0.00	1909	2	J		29	0.00	219	65	45	0		Revf	Sist	
104	AL VOERTE	39.10	20.40	12.0	3.0	0.3	- 4	0.0	0.00	1909	3	23	13	22	0.00	80	40	104	0		Norf		WAE
105	KADABUDUN	30.33	20.40	4.0	3.9	0.3	8	0.0	0.00	1969	3	26	1	48	0.00	281	34	-90	36	13	Norf Norf Norf		WAI
106	CERTS	30.41	20.41	10.0	3.0	3.6		0.0	0.00	1909	4	. 6	_Z	49	0.00	116			0	0	Norf		WAI
107	VIII AUTO	39.21	29.31	10.0	0.0	7 - 1	9	0.0	0.00	1970	3	28	21	2	0.00	281			40	225	Norf		WAE
107	NUTARYA	39.10	29.80	26.0	5.5	0.0	В	0.0	0.QC	1970	4	19	13	47	0.00	0	0						WAI
	DEMIRCI-GORDES	39.10	28.65	28.0	5.2	0.0	В.	0.0	0.00	1970	4	23	9	1	0.00	0			0	0			WA
109	GONEN DEMIRCI - GORDES ALASEHIR KARABURUN GEDIZ KUTAHYA DEMIRCI - GORDES GURUN - SIVAS IVRINDI - BALIKESIR RUEDUR	38.87	36.81	19.0	4.7	+ 8	7	0.0	0.00	1970	7	2	2	24	0.00	193	92	- 33	a	0	Slst		CAI
110	IVRINDI-BALIKESIK	39.62	27.32	10.0	5.0	5 6	9	0.0	0.00	1971	2	23	19	41	0.00	0	0	0	0	0			WA
111	BURDUR	37.60	29.72	30.0	6.0	6 2	8	0.0	0.00	197 L	5	12	6	25	0.00	64	50	-75	40	30	NOFT		WAI
112	BINGOL	38.85	40.52	9.0	5.9	5.7	8	0.0	0.00	1971	5	22	16	43	0.00	231	81	4	38	20	Sist		EA
113	SARIKAMIS	40.42	43.88	46.0	4.9	0.0	8	0.0	0.00	1972	3	22	0	51	0.00	0	0	0	0	0			NE
114	VAN-EDREMIT	38.23	43.36	46.0	4.9	0.0	8	0.0	0.00	1972	7	16	2	46	0.00	0	0	Q	0	0			EA
115	IZMIR	38.55	27.22	24.0	5.2	5.8	7	0.0	0.00	1974	2	1	0	1	0.00	250	65	151	0	à	Norf		WA
116	GELIBOLU	40.45	26.12	15.0	3.5	5.6	7	0.0	0.00	1975	3	27	5	15	0.00	279	46	-43	Ó	Ô	Norf		NA
117	LICE	38.51	40.77	32.0	6.0	6.6	В	0.0	0.00	1975	9	6	9	20	0.00	250	54	43	26	60	Revf	Slat	EA
116	ARDAHAN	41.01	42.97	25.0	4.8	5.1	. 8	0.0	0.00	1976	3	25	11	55	0.00	0	0	0	0	0	•		NE.
119	IGDIR	39.85	43.64	14.0	4.6	5 1	7	0.0	0.00	1976	4	2	16	58	0.00	o	o	0	ā	ā			NE
120	ARDAHAN	40.96	42.87	30.0	5.0	5.5	8	0.0	0.00	1976	4	29	22	18	0.00	ō	ō	ō	ã	ã			NE
121	DENIZLI	37.71	29.00	20.0	5.0	4.9	6	340.0	0.50	1976	8	19	1	12	0.60	ō	ō	ō	ō	ō	Norf		WAI
122	CALDIRAN	39.05	44.00	10.0	6.9	7.3	9	335.0	0.00	1976	11	24	12	22	0.71	196	86	166	54	370	Sler		WAL
123	PALU-ELAZIG	38.58	40.03	29.0	5.0	4.8	7	0.0	0.00	1977	3	25		39	0.00	232	86	-8	'n	ň	UISC		PAG
124	IZMIR	38.41	27.19	24.0	5.3	5 6	7	221.0	0.50	1977	12	16	7	37	0.19		ñ	ň	ň	ň	Nort		WAS
125	UZUMLU-ERZINCAN	39.67	39.88	48.0	4.7	4.5	7	0.0	0.00	1978	2	15	3	17	0.00	ō	ō	ñ	ñ	ŏ			MAI
126	BURDUR	37.53	30.02	28.0	4.6	4.5	7	126.0	0.00	1978	7	29	64	34	0.00	Ď	0	ñ	ň	ň			WAS
127	KARABURUN	38.79	26.57	15.0	5.9	5.8	6	137.0	0.00	1979	6	14	11	44	0.00	õ	ŏ	ñ	ň	ŏ			WAI
128	PULUMUR-TUNCELI	40.65	42.75	3.0	4.3	0.0	7	0.0	0.00	1980	4	25	-ō	57	0.00	ñ	ñ	ñ	ň	ň			NA.
129	TERCAN-ERZURUM	39.90	40.31	37.0	5.L	4 4	7	0.0	0.00	1980	11	18	3	14	0.00	õ	ñ	ň	ň	ň			ME
130	BULANIK-MUS	39.23	41.90	38.0	5.4	5.2	7	0.0	0.00	1982	3	27	19	57	0.00	õ	ñ	ŏ	ň	Ö			NAI
131	HORASAN	40.07	42.26	62.0	4.7	3.8	7	0.0	0.00	1982	5	19	13	32	0.00	ñ	ñ	ŏ	ñ	ň			NE
132	CAY(RLI-TERCAN	39.89	40.43	45.0	5.0	4.1	7	0.0	0.00	1983	4	6	7	35	0.00	n	ň	ñ	ň	ň			NE
133	KULU-BALA	39.30	33.06	36.0	4.7	4.4	7	0.0	0.00	1983	4	21	16	18	0.00	ñ	0	ñ	ñ	ň			Cal
134	BIGA	40.33	27.21	7.0	5.5	5.8	7	51.0	0.00	1983	7	- 5	12	1	0.68	4K	50	a ₁	ñ	ň	Pauf		GA.
135	INEGOL-BURSA	40.14	29.35	12.0	5.0	4.9	7	0.0	0.00	1983	10	21	20	3.3	0.00	ñ	ñ	ň	ň	ň	KUVI		NA
136	URUN-SIVAS BURDUR BINOOL SARIKAMIS VAN-EDREMIT IZMIR OGLIBOLU LICUE ARDAHAN IGOLE ARDA	40.35	42.18	9.0	6.0	6.8	8	180.0	0.80	1983	10	30	4	12	0.42	215	64	7	Ř	100	Slat		NP
137	BALKAYA-ERZURUM	40.90	42.24	10.0	5.3	5.5	7	0.0	0.00	1984	9	18	13	26	0.00	112	79	71	0	-00	Sigr		NP
138	DOCANSEHIR-MALATYA	38.02	37.79	10.0	5.7	5.8	7	56.0	0.00	1986	6	6	3	35	0 33	57	g	60	20	ň	Revf		EAL
	DOGANSEHIR-MALATYA KUYUCAK-DENIZLI NARMARA YURADIYE DURSUNBEY ANTAKYA BALIKESIR CAYURLI-ERZINCAN KOYCEGIZ ERZINCAN PULUMUR DINAR	37.94	28.56	5.0	5.4	5.7	5	0.0	0.00	1986	10	11	ğ	ő	0.00	ā	ñ	ñ	20	ä	~~ v i		WAI
140	MARMARA	40.93	28.07	11.0	5.0	5.1	7	0.0	0.00	1988	- 4	žì	20	40	0.00	ă	ñ	ñ	ŏ	a			NAE
141	MURADIYE	39.12	43.91	44.0	4.9	5.1	Ó	49.8	0.20	1979	4	11	-0	0	0.48	á	n	ň	ñ		Sist		EAF
142	DURSUNBEY	39.66	28.65	7.0	5.2	5.4	ō	309.4	0.40	1979	÷	18	10	13	0.52	ñ	ň	ň	a		Norf		WAI
143	ANTAKYA	36.17	35.89	63.0	4.7	4.8	6	128.8	0.12	1981	6	30	ĩă	59	0 97	ň	ň	ň	ă		MOLI		EAF
144	BALIKESIR	39.64	28.87	12.0	4.6	4.7	ŏ	295.7	0.35	1984	3	20	ໍ້າ	6	1 22	ă	0	0	0		Norf		WAF
145	CAYURLI-ERZINCAN	39.95	39.77	29.0	4.9	5.0	ŏ	152.4	0.50	1985	Ä	12	2	4.1	0.28	0	ď	0	0	0	not t		
146	KOYCEGIZ	37.00	28.70	0.0	3.5	3.6	ő	144.7	0.67	1985	12	7	â	77	0.48	ŏ	0	Ň	0	0			NAL
147	ERZINCAN	39.71	39.55	29.0	6.2	6.8	ă	505.0	0.50	1007	* *	12	3	10	0.40	101	D. 1	70	45		V	01	WAE
148	PULUMUR	39.53	39.87	16.0	5.5	6.1	7	100.8	0.50	1992	3	15	18	17	0.30	61	70	1.4	45	94	Norf	Sist	NAE
149	DINAR	38.10	30.02	24.0	5.7	6.0	ò	294.0	0.00	1995	10	+ 5		57	0.91	91	10	-4	-0		W#	01-4	NAE
150	DINAR	38.10	79 87	20.0	0.0	1.7	ŏ	174 0	0.00	1005	10	26	13	J/	0.38	Ü	ú	Ü	50			Slst	WAF
	- a 111	30.10	20.04	40.0	4.0	4./	U	1/4.U	0.00	TAAO	9	60	1.4			0	0	D	0		Norf	Sist	WAE

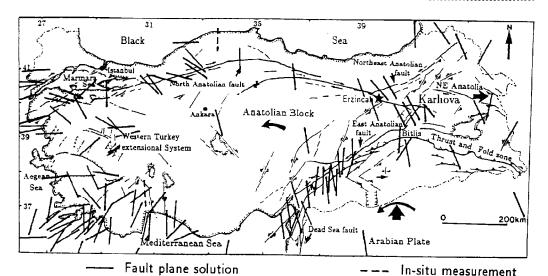


Fig. 1 Main fault systems and Tectonics of Turkey (thick linear solid and broken lines show the maximum compressive horizontal stress directions)

System for Interrelation Studies

A soft-wear program coded with BASIC programing language was developed to create data files for studying interrelations among seismic characteristics of earthquake data. This soft-wear is presently capable of creating data files for the following conditions:

- 1) Rupture type (Slst, Norf, Revf)
- 2) Earthquake grouping (NAF, NEAF, WAFS, CAFS, EAFBZ)
- 3) Magnitude range (M_{min}, M_{max})
- 4) Unconditional

Created data files were then plotted using a graphic softwear NGRAPH, examples of which are presented in the following section.

SEISMIC CHARACTERISTICS AND INTERRELATIONS

Locations of Earthquakes for the Last Century

Figure 2 shows plots of locations of earthquakes in Turkey for the last century which corresponds to the instrumented period. As seen from the figure, earthquakes concentrate along the North Anatolian Fault (NAF), North East Anatolian Fault (NEAF), Central Anatolian Fault System (CAFS), East Anatolian Fault and Bitlis-Zagros Thrust Zone (EAFBZ), West Anatolian Fault System (WAFS) as expected.

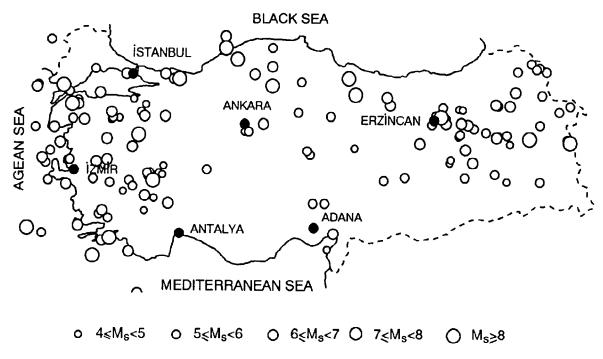


Fig. 2 Locations of earthquakes in Turkey for the last century

Relation between Magnitude M_s and Mercalli Intensity I_o

Figure 3 shows the relation between magnitude M_s and Mercalli Intensity I_o . The following linear functions were fitted to observations.

$$I_o = 1.317 M_s, \quad M_s = 0.7593 I_o$$
 (1)

It should be noted a constraint was implemented in fitting the emprical equation (1) to observations.

Relation between Magnitude M, and Maximum Displacement U_{max}

Figure 4 shows the relation between magnitude M_s and maximum observed relative displacement U_{max} . As seen from the figure, the observation of movement of faults on ground surface becomes possible when the magnitude M_s is greater than 4. The following emprical function seems to be a good fit to observations.

$$U_{max} = 0.0014525 M_s e^{1.31 M_s} (2)$$

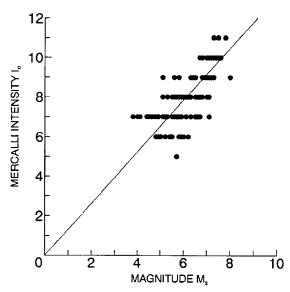


Fig. 3 Relation between magnitude M_s and Mercalli Intensity I_o

As expected, the magnitude of displacement of the fault increases as the magnitude of the earthquake increases.

Relation between Magnitude M, and Surface Rupture Length L

Figure 5 shows the relation between magnitude M_s and surface rupture length L of faults. Eq. (2) exactly fits the observations except that coefficient 0.0014525 has a different physical meaning. This implies that the surface rupture length is 10^5 times maximum relative displacement U_{max} for Turkish earthquakes which is not very suprising if the definition of the seismic moment is considered.

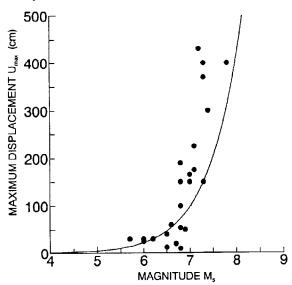


Fig. 4 Relation between magnitude M_s and maximum displacement U_{max}

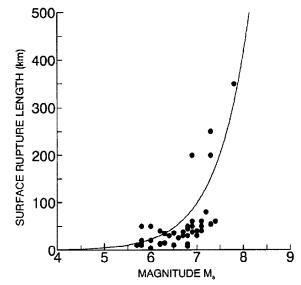


Fig. 5 Relation between magnitude M_s and surface rupture length L

Relation between Magnitude M_s and Rupture Area A

Figure 6 shows the relation between magnitude M_s and rupture area A. The following emprical function seems to be representative of observations.

$$A = 0.0032 M_s \epsilon^{1.66 M_s} \tag{3}$$

Relation between Magnitude M_s and Peak Ground Acceleration a_{max}

Figure 7 shows the relation between magnitude M_s and peak ground acceleration. It should be noted that hypocenter distance R from strong motion stations was generally greater than 20 km. The following emprical function seems to be a good fit to observations.

$$a_{max} = 2.8(e^{0.9M_{\bullet}}e^{-0.025R} - 1)$$
 (4)

By assuming that R=25km, the plot of the above function is also shown in the same figure. This function is a convolution of two exponential functions and its coefficients may be further refined as more strong motion data increases.

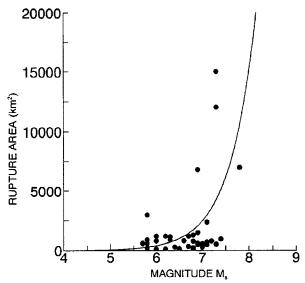


Fig. 6 Relation between magnitude M_s and rupture area A

Relation between Magnitude M_s and Dominant Period τ_{per}

Figure 8 shows the relation between magnitude M_s and dominant period τ_{per} of observed acceleration waves. The dominant period is the one that corresponds to the spectrum of the wave with the maximum amplitude in Fourier Spectra. The following emprical function seems to be representative of observations.

$$\tau_{per} = 0.067 + 0.071M_s \tag{5}$$

Although the number of data is not sufficient to draw any conclusion yet, the period of the dominant wave should generally be between 0.2-0.8 second in view of fracture propagation velocities reported in rock mechanics.

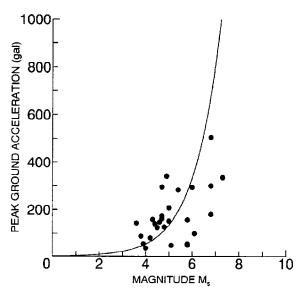


Fig. 7 Relation between magnitude M_s and maximum acceleration PGA

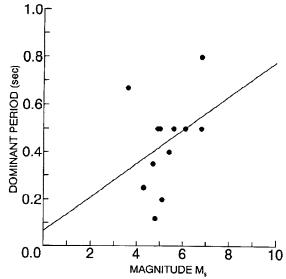


Fig. 8 Relation between magnitude M_s and dominant period of acceleration wave τ_{per}

Relation between Magnitude M_s and Ratio of Maximum Vertical Acceleration to Maximum Horizontal Acceleration RVAHA

Figure 9 shows the relation between magnitude M_s and the ratio of maximum vertical acceleration to maximum horizontal acceleration RVAHA. The following emprical function fits the observations.

$$RVAHA = 0.217 + 0.046M_s \tag{6}$$

Since most of earthquakes are of in-land type, the ratio ranges between 0.2 to 0.8. However, the average seems to be 0.4 which implies importance of considering vertical component in the seismic design of engineering structures against in-land earthquakes.

Attenuation of Peak Ground Acceleration

Figure 10 shows the relation between hypocenter distance R and peak ground acceleration a_{max} together with a function given by Eq. (4), which is plotted for various values of magnitude M_s . It should be noted that the data plotted in Fig. 10 were obtained from another data-base system named IVAZTURD for the attenuation and wave characteristics of acceleration records for the strong motion wave data of Turkey. Since the number of strong motion data for $Turkish\ Earthquakes$ is still limited, it is in-mature to draw any conclusion yet. Nevertheless, the proposed function presently seems to be a good-fit to observations.

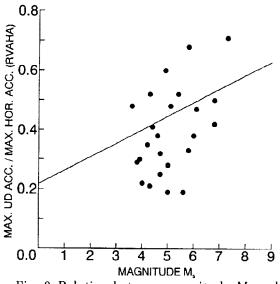


Fig. 9 Relation between magnitude M_s and maximum vertical acceleration to maximum horizontal acceleration

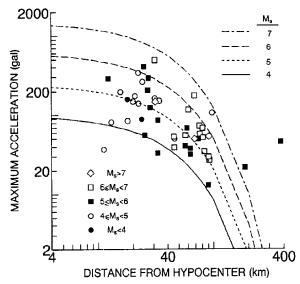


Fig. 10 Attenuation of maximum ground acceleration

EARTHQUAKE OCCURENCE PATTERN IN TURKEY

Figure 11 shows the relation between time (year) and Magnitude M_s of earthquakes for various faulting groups of Turkey. As it is noted from the figure, very large events $(M_s \ge 7)$ occur at the NAF and at the EAFBZ while the magnitude of events at the WAFS and the NEAF are generally less than M_s =7. It seems that the initiation of seismic activity occurs in the near vicinity of Karhova junction (see Fig. 1 for location) in the form of lateral strike slip and/or reverse faulting. Reverse faulting events are generally shallow (h < 15km) while strike-slip faulting events are more deeper. As a result, events are simultaneously or with a slight delay initiated in the WAFS in the form of normal faulting. For example, the Kuşadası earthquake $(M_L = 6.0)$ on 06/11/1992 and the recent Dinar earthquake $(M_L = 6.0)$ on 01/10/1995 followed the Erzincan earthquake $(M_L = 6.9)$ on 13/03/1995 and Pülümür earthquake $(M_L = 6.0)$ on 15/03/1995 confirm this conclusion. The events in the WAFS have smaller magnitudes as the stress state to cause normal faulting could not result in higher energy accumulation as compared with those in the NAF and the EAF. However, the frequency of events in the WAFS is greater than that of the eastern events. This is mechanically a quite natural consequence of the stress states to cause normal faulting and strike-slip faulting.

Although it is still inmature to purpose an exact model for earthquake occurrence pattern in Turkey, the following model may be put forward in view of the available data:

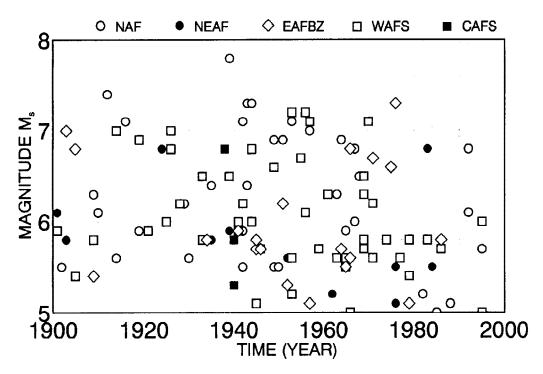


Fig. 11 Distribution of earthquakes in time

- 1) Crushing and strike-slip faulting in the vicinity of the Karhova junction of the NAF, NEAF and EAFBZ: The driving force results from the northward moving and anti-clock wisely rotating Arabian plate. It must be noted that it is almost imposible to push the Anatolian plate westward by a purely northward-moving Arabian plate. If the anti-clock wise rotation of the Arabian plate exists, then it will provide an enough momentum to push the Anatolian plate towards the west. Recent Çaldıran earthquake (24/11/1976) validates the above statement.
- 2) Uplifting and normal faulting in the western Anatolia: The westwardly pushed Anatolian plate is restrained by the Agean plate and is uplifted by the northwardly moving African continent against the stationary Euro-Asian plate. This results in bending the western part of the Anatolian plate which causes normal faulting in the uppermost part of the crust. This results in a further growth of the host-graben structure in the western Anatolia. However, it should be noted that there is a general tendency for uplifting the whole western Anatolia. For example, the historic sea-side cities such as Efes (Ephesos) and Milet (Miletos) in the grabens of Küçük Menderes and Büyük Menderes respectively are retreated by 3 to 4 km away from the present sea shore, which validates the above conclusion.
- 3) Rotation of the Anatolian Plate: Because of the combined action of the slower rate of motion of the African plate with respect to that of the Arabian plate, westward pushing by the Arabian plate, restraining by the Agean plate and uplifting by the African plate, the Anatolian plate tends to rotate anti-clock wise, aided with the motion of the rotating Earth. This action causes subsequent events along the entire perimetry of the Anatolian plate. Without doubt the magnitude of events at bends, where stress concentrations occur, are larger than at other less restraining segments of the faults.

The information from the present data-base system and previous two dimensional elasto-plastic finite element analyses of Turkey and its close vicinity (Kasapoğlu and Toksöz 1983) confirms most of the above statements. Nevertheless, three-dimensional elasto-visco-plastic analyses with the consideration of the actual geologic structure of the Earth's crust in Turkey and its close vicinity are necessary. If these analyses are backed up by in-situ stress measurements and global and local deformation measurements, they can probably yield an exact model for the earthquake occurrence pattern in Turkey.

CONCLUSIONS

An integrated data-base system TURDEP for Turkish Earthquakes was developed to study the seismic characteristics of the earthquakes of Turkey so that it may be of some help to engineers for the seismic design of structures as well as to scientists for forecasting earthquakes. This data-base system was described and the usefullness of the system was demonstrated by studying the interrelations among the seismic characteristics of the Turkish earthquakes, inspite the system still needs further improvements. It is also pointed out that such systems integrated with numerical analyses and monitoring schemes can be of great value in forecasting future earthquakes not only in Turkey but also in other countries prone to earthquakes.

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